

QUANTIFYING AND USING UNCERTAINTY: REVIEW AND HIGHLIGHTS

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Uncertainty: meaning and importance

Uncertainty answers this question:

Given a value, within what spread of values is it reasonable to consider that the truth lay?





STATION DATA

http://www.ecad.eu/

See also Antonello Squintu's poster!



Many of these series have inhomogeneities







BREAK DETECTION

Annual mean original tn 127 Bamberg GERMANY





THE ADJUSTMENTS

Adj. est. final1, month 5, ser id 127, break 1952

Each coloured line gives the adjustment against a different neighbouring station







EUSTACE and Uncertainty

Integrated approach across project

Common understanding & models of uncertainty

Outline:

- Introductory comments (done that!)
- Station Data (done that!)
- Satelltie ST retrieval and uncertainty validation
- L2 to L3 uncertainty propagation
- Satellite-to- T_{2m}
- Usage of uncertainty information in SAT analysis



Sources and propagation of uncertainty





Sources and propagation of uncertainty





Satellite Surface Temperature Data

- Common three-component uncertainty model
 - random
 - locally correlated
 - systematic
- Validation of uncertainty
- Three-component model
 - Applies to all domains, land, ice, lake, sea
 - Applies across processing levels
 - Provides information propagated into analysis



Land ST Uncertainty Components

VARIABLE	Метнор	Comments	
LST_UNC_RAN	L2 Random 1 / Radiance noise Propagation	$u_{ran,y}(x) = \sqrt{\sum_{c=1}^{n} \left(\frac{\partial R}{\partial y_c} u_{ran}(y_c)\right)^2}$	Random component of L1 channel uncertainties propagated through the retrieval
	L2 Random 2 / Emissivity noise Propagation	$u_{ran,\varepsilon}(x) = \sqrt{\sum_{c=1}^{n} \left(\frac{\partial R}{\partial \varepsilon_{c}} u_{ran}(\varepsilon_{c})\right)^{2}}$	Estimate of the magnitude of pixel-to-pixel scale emissivity variability within areas based on land cover class
LST_UNC_LOC	L2 Local 2 / Uncertainty from atmosphere/fit for regression-based retrieval L2 Local 2 / Uncertainty from Emissivity	$u_{loc,fit}(x) = \sqrt{Var(\hat{x} - x_{in})}$ $u_{loc,\varepsilon}(x) = \sqrt{\sum_{c=1}^{n} \left(\frac{\partial R}{\partial \varepsilon_c} u_{loc}(\varepsilon_c)\right)^2}$	Atmospheric fields correlated on timescales >1 day and length scales >100 km. For coefficient based retrieval methods the retrieval ambiguity is a contributor of residuals in the fit Across a particular land class area, there may be a mean difference between the assumed and true mean emissivity
LST_UNC_SYS	L2 Systematic 1 / Reasoned estimate	Assumed that known corrections have been applied by data producers and what remains is describable as an uncertainty in the bias of the satellite surface temperatures relative to other data sources of temperature (ie from validation)	





National Centre for Earth Observation



MODIS LST Uncertainties





Locally correlated - atmosphere





National Centre for Earth Observation



Locally correlated - surface



SEVIRI LST Uncertainties

Random

Locally correlated



LST_unc_ran [K]

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LST_unc_loc [K]







VALIDATION OF SATELLITE UNCERTAINTIES

• Test the goodness-of-fit between the uncertainty from in situ validation ($\sigma_{sat-ground}$) and the total satellite product uncertainty for each associated matchup (σ_{total})

•
$$\sigma_{total} = \sqrt{\sigma_{sat}^2 + \sigma_{ground}^2 + \sigma_{space}^2 + \sigma_{time}^2 (+ \sigma_{depth}^2)}$$

- σ_{sat} is the total LST uncertainty for satellite pixel
- σ_{ground} is the uncertainty associated with the ground-based measurement
- σ_{space} is the uncertainty associated with matching a satellite and ground observation in a spatial context
- σ_{time} is the uncertainty associated with matching a satellite and ground observation in time
- σ_{depth} is the uncertainty due to the difference in depth of the measurements (SST only)







IST UNCERTAINTY VERIFICATION

• Validation against independent radiometer observations from ARMS



AASTI IST uncertainty validation with respect to ARM in-situ data for 2008. Dashed lines show ideal uncertainty model accounting for uncertainties in the in situ data and geophysical uncertainties arising from spatial and temporal collocation. Solid black lines show one standard deviation of the retrieved minus in situ IST differences for each 0.1 K bin.



Sources and propagation of uncertainty





PROPAGATION OF L2 -> L3 UNCERTAINTIES

LST: Uncertainties propagated from 1km LST pixels (*upixel*) -> global 0.25 grid LST (*ucell*):







Sources and propagation of uncertainty





SIMPLE AIR-SEA TEMPERATURE DIFFERENCE MODEL

THE THING DESIRED Climatological Offset – Fourier components

MAT =
$$SST + \delta + \varepsilon$$

Measured by ships only

Measured by satellite/ship

Temporally and spatially correlated variability





Data from ICOADS 2.5 1963-2000



UNCERTAINTY IN SATELLITE LSAT

• Simple multiple linear regression model to estimate LSAT:

$$T_{max} = \alpha_{0} + \alpha_{1} . LST_{day} + \alpha_{2} . LST_{ngt} + \alpha_{3} . FVC + \alpha_{4} . SZA_{noon} + \alpha_{5} . Snow + \varepsilon_{Tmax}$$

$$T_{min} = \beta_{0} + \beta_{1} . LST_{day} + \beta_{2} . LST_{ngt} + \beta_{3} . FVC + \beta_{4} . SZA_{noon} + \beta_{5} . Snow + \varepsilon_{Tmin}$$

$$Vegetation \uparrow Solar Zenith angle at noon from fraction fracti$$

Surface

Met Office Hadley Centre

Ra

$$Tmax_{surf} = (\alpha_1^2 . LST_{day_ucell_surf}^2 + \alpha_2^2 . LST_{ngt_ucell_surf}^2 + \alpha_3^2 . FVC_{ucell_local}^2)^{\frac{1}{2}}$$
$$Tmin_{surf} = (\beta_1^2 . LST_{day_ucell_surf}^2 + \beta_2^2 . LST_{ngt_ucell_surf}^2 + \beta_3^2 . FVC_{ucell_local}^2)^{\frac{1}{2}}$$
$$EUSTACE$$

EXAMPLE UNCERTAINTY FIELDS (1 JULY 2010)



Sources and propagation of uncertainty



Advanced Standard Air Temperature Model

Temperature Process Decomposition

Temperature variability is decomposed into model sub-components with defined structure in space/time:

$$T(s,t) = T^{\text{clim}}(s,t) + T^{\text{large}}(s,t) + T^{\text{local}}(s,t)$$

T(s,t) = Temperature at space/time location (s,t) $T^{\rm clim}(s,t)$ = Climatological temperature $T^{\rm large}(s,t)$ = Large spatial/temporal scale component $T^{\rm local}(s,t)$ = Daily, short spatial scale component





Temperature Observation Model

Observation model

Daily mean air temperatures are decomposed into variability at different scales:

$$y^i = T(s^i, t^i) + b^i + \epsilon^i$$

Where b^i is a sum of observational biases affecting observation i and ϵ^i are non-bias related observational errors.







Analysis method

Met Office

Spatial interpolation based on the SPDE approach (Lindgren et al 2011):

- Temperatures are modelled as weighted sum of local functions.
- A Bayesian method, where variability/smoothness is controlled by a prior distribution for the weights.
- Compute the probability density function of the weights conditioned on the temperature observations.



Lindgren, F., H. Rue, J. Lindström, (2011). An explicit link between Gaussian fields and Gaussian Varkov random fields: the stochastic partial differential equation approach, *Journal of the Royal* Statistical Society: Series B (Statistical Methodology), 73, 4



Analysis method

Estimation of temperatures and observation biases

Jointly estimate temperature model variables u and observation bias variables eta, with $x = (u, \beta)$ and observations:

$$egin{aligned} &y = {J}_{oldsymbol{u}} oldsymbol{u} + {J}_{eta}eta + \epsilon \ &= {J}_{oldsymbol{x}} oldsymbol{x} + \epsilon \end{aligned}$$

Apply Bayes' Rule to compute the analysis:

 $\begin{array}{ll} \text{Prior}: & (\boldsymbol{x} \mid \boldsymbol{\theta}) \sim \mathcal{N}(\boldsymbol{\mu}_{\boldsymbol{x}}, \boldsymbol{Q}_{\boldsymbol{x}}^{-1})\\ \text{Observations}: & (\boldsymbol{y} \mid \boldsymbol{x}, \boldsymbol{\theta}) \sim \mathcal{N}(\boldsymbol{x}, \boldsymbol{Q}_{\boldsymbol{\epsilon}}^{-1})\\ \text{Posterior}: & p(\boldsymbol{x} \mid \boldsymbol{y}, \boldsymbol{\theta}) \propto p(\boldsymbol{y} \mid \boldsymbol{x}, \boldsymbol{\theta}) p(\boldsymbol{x} \mid \boldsymbol{\theta}) \end{array}$

Observational uncertainties are encoded into Q_{ϵ} . Structured errors have their structure encoded into J_{β} and their magnitude β is estimated.





Model Component Solutions

Linear Model

Each of the three model components are constructed as a linear (or linearised) model, with a design matrix J and latent variables to be estimated x:

$$y = Jx + \epsilon$$

- Measurement error is additive Gaussian $p(\epsilon) = \mathcal{N}(\mathbf{0}, \boldsymbol{Q}_{\epsilon}^{-1})$
- Model variables x have a Gaussian prior distribution $p(x) = \mathcal{N}(\mu_x, Q_x^{-1})$

Estimation

Met Office

Compute the distribution of \boldsymbol{x} conditioned on the observations:

$$\begin{split} p(\boldsymbol{x} \mid \boldsymbol{y}, \boldsymbol{\theta} = \hat{\boldsymbol{\theta}}) &= \mathcal{N}(\boldsymbol{\mu}_{\boldsymbol{x}\mid\boldsymbol{y}}, \boldsymbol{Q}_{\boldsymbol{x}\mid\boldsymbol{y}}^{-1}) \\ \boldsymbol{\mu}_{\boldsymbol{x}\mid\boldsymbol{y}} &= \boldsymbol{\mu}_{\boldsymbol{x}} + \boldsymbol{Q}_{\boldsymbol{x}\mid\boldsymbol{y}}^{-1} \boldsymbol{J}^T \boldsymbol{Q}_{\boldsymbol{\epsilon}} (\boldsymbol{y} - \boldsymbol{J} \boldsymbol{\mu}_{\boldsymbol{x}}) \\ \boldsymbol{Q}_{\boldsymbol{x}\mid\boldsymbol{y}} &= \boldsymbol{Q}_{\boldsymbol{x}} + \boldsymbol{J}^T \boldsymbol{Q}_{\boldsymbol{\epsilon}} \boldsymbol{J} \end{split}$$

Efficient solution depends on the sparse structure of J , Q_x and $Q_\epsilon.$



Demonstration Application

- Demonstrated on small region/subset of input data.
- Applied to in situ air temperature (HadNMAT2, GHCN Daily) and satellite LST derived air temperature.
- Climatology fitted using observations in 1961-1990.
- Placeholder uncertainty information.







Daily local uncertainty







CONCLUSIONS



- EUSTACE attempting an integrated and coherent treatment of uncertainty at all levels of data
- Input data uncertainties have been estimated and validated
- Propagated and introduced uncertainty characterised as required at each step

 L1→L2→L3→Analysis
- Coherency across observational epochs through consistent statistical treatment of uncertainty



